

# ANALYSIS OF SOME NONLINEAR PDES FROM MULTI-SCALE GEOPHYSICAL APPLICATIONS

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“Mathematics for the Fluid Earth”

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06 November, 2013

## Objects

**Multi-time-scale** PDE systems from geophysical applications, solutions fast oscillating in  $t$ . Nonlinearity generates resonances.

$$\partial_t U = \varepsilon^{-1} \mathcal{L}[U] + F$$

- Rotating shallow water eq in  $\Omega \subset \mathbf{R}^2$ ; incompressible fluid dynamics on fast rotating  $\mathbf{S}^2$ .

## Objectives

**Estimate**  $\mathcal{P}(\text{multiscale soln}) - (\text{limiting soln})$  in terms of  $\varepsilon \ll 1$ .

- Multiscale solution is decomposed by  $\mathcal{P}$ , into slow component in  $\ker\{\mathcal{L}\}$  and fast component.
- Initial data are **ill-prepared** (unfiltered), i.e.  $O(1)$  fast component.

## Methodology

- **Time-averaging** to suppress fast component.
- Estimate resonances using **vorticity equation**

vorticity dynamics  $\equiv$  slow dynamics

### Relation to numerical analysis

- Numerical methods can be validated by showing that they give the theoretically predicted rate of convergence to limiting solutions.

$\|\mathcal{P}u_{\Delta x}^\varepsilon - u_{\Delta x}^0\| \sim O(\varepsilon + (\Delta x)^n)$ . Then,

$$\|\mathcal{P}(u_{\Delta x}^\varepsilon - u^\varepsilon)\| \leq \|\mathcal{P}u_{\Delta x}^\varepsilon - u_{\Delta x}^0\| + \|u_{\Delta x}^0 - u^0\| + \underbrace{\|u^0 - \mathcal{P}u^\varepsilon\|}_{\text{this talk}}$$

$u_{\Delta x}^\varepsilon$ : multiscale numerical soln.

$u^\varepsilon$ : multiscale exact soln.

$u_{\Delta x}^0$ : limiting numerical soln.

$u^0$ : limiting exact soln.

(Direct error estimates:  $\|\mathcal{P}u_{\Delta x}^\varepsilon - \mathcal{P}u^\varepsilon\| \sim (\Delta x)^m/\varepsilon$ , not good if  $\varepsilon \ll 1$ .)

Multi-scale-ness is measured by parameters. As some parameters  $\rightarrow 0/\infty$ , reduced models *may* still capture relevant features.

Rotating shallow water (RSW) eq. as **singular perturbation** problem

$$\left\{ \begin{array}{l} \partial_t H + \nabla \cdot (Hu) = 0, \\ \partial_t u + u \cdot \nabla u + \frac{\nabla H}{Fr^2} + \frac{u^\perp}{Ro} = 0, \end{array} \right. \quad \text{subject to } u \cdot \mathbf{n} \Big|_{\partial\Omega} = 0$$

$H$  is total height. Parameters: 2D **Froude** and **Rossby** numbers.

With  $Fr = Ro = \varepsilon$ ,  $H = 1 + \varepsilon h$ , system fits  $\partial_t U = \varepsilon^{-1} \mathcal{L}[U] + F$ .

**But**,  $Fr$  and  $Ro$  can scale differently.

- $Fr^2 = Ro \ll 1$ : semi-geostrophic (SG) [Phillips 63'; Brenier, Cullen, Feldman, Roulstone, Oliver, etc etc 00s]
- $Fr^2 \ll Ro \ll 1$ : quasi-geostrophic (QG) [Cheng 08', 14'].
- $Fr^2 \gg Ro$ : near-inertia oscillation [Cheng & Tadmor 08']

## Well-prepared vs ill-prepared data

From here on, use  $U$  for  $\varepsilon$ -dependent solution,  $\tilde{U}$  for limiting soln.

$$\varepsilon(\partial_t U + \text{nonlinear}) + \mathcal{L}[U] = 0.$$

Always assume  $U$  is smooth for  $t \lesssim 1$  with spatial norms  $\sim O(1)$ .

	<b>Well-prepared</b>	<b>Ill-prepared</b>
<i>Initial conditions</i>	<b>Impose</b> $\partial_t U_0 \sim O(1)$ iff $\mathcal{L}[U_0] \sim O(\varepsilon)$	<b>Allow</b> $\partial_t U_0 \sim O(\varepsilon^{-1})$ iff $\mathcal{L}[U_0] \sim O(1)$
<i>For local times <math>t</math></i>	Above persists	Above persists if w/o dispersion/dissipation
$\tilde{U} \stackrel{?}{=} \lim_{\varepsilon \rightarrow 0} U$	Possible	Trouble!

System of slow/fast scales

$$\partial_t U = \varepsilon^{-1} \mathcal{L}[U] + F$$

$\mathcal{L}$ : linear operator with purely imaginary spectrum (incl. 0).

$F$ : nonlinear and source terms, with spatial norms  $\sim O(1)$ .

## Lorentz

$\mathcal{P}$  : some projection onto  $\ker\{\mathcal{L}\}$ ,  $U^s := \mathcal{P}U$ ,  $U^f := U - U^s$ .

The solution space is decomposed into a slow subspace  $\ker\{\mathcal{L}\} \ni U^s$ , and a fast subspace  $\ni U^f$ .

- $\partial_t U^s \sim O(1)$ , so it is possible to prove  $\exists \lim_{\varepsilon \rightarrow 0} U^s$  by Arzelà-Ascoli
- $\ker\{\mathcal{L}\}$  is in general not invariant w.r.t. original, nonlinear dynamics.

For  $U = \begin{pmatrix} h \\ u \end{pmatrix}$ ,  $\mathcal{L}[U] = \begin{pmatrix} \nabla \cdot u \\ \nabla h + u^\perp \end{pmatrix}$ . So,  $\mathcal{L}[U] = 0$  iff  $u = \nabla^\perp h$

$$U^s = \mathcal{P}U := \begin{pmatrix} \psi \\ \nabla^\perp \psi \end{pmatrix} \text{ where } \psi := (\Delta - 1)_{\text{QG}}^{-1}(\nabla \times u - h)$$

- $(\Delta - 1)_{\text{QG}}^{-1} \psi$  enforces  $\psi|_{\partial\Omega} = \text{constant}$ .
- $(\nabla \times u - h)$  is linear potential vorticity;  $\mathcal{P}$  is linear *PV inversion*, analogue of Biot-Savart law.

For the “limiting” solution  $\tilde{U} = \lim_{\varepsilon \rightarrow 0} U^s$

WKB expansion in  $\varepsilon$  with *well-prepared* data  $\implies$

- Leading order. No dynamics;  $\tilde{U} \in \ker\{\mathcal{L}\}$ , i.e.  $\tilde{u} = \nabla^\perp \tilde{h}$
- Next order. Dynamics; 2D **quasi-geostrophic (QG) eq.**

2D RSW eq. with  $u \cdot \mathbf{n}|_{\partial\Omega} = 0$ , (fast+slow scales,  $U = U^s + U^f$ )

$$\begin{cases} \partial_t h + \nabla \cdot (hu) + \varepsilon^{-1} \nabla \cdot u = 0, \\ \partial_t u + u \cdot \nabla u + \varepsilon^{-1} u^\perp + \varepsilon^{-1} \nabla h = 0, \end{cases} \quad (h_0, u_0) \text{ given.}$$

2D QG eq., (slow scale,  $\tilde{U} = \mathcal{P}\tilde{U} = \tilde{U}^s$ )

$$\begin{cases} \partial_t \tilde{\omega} + \tilde{u} \cdot \nabla \tilde{\omega} = 0 & \tilde{\omega}_0 = \nabla \times u_0 - h_0 & (\tilde{\omega} \text{ is PV}) \\ \tilde{h} = (\Delta - 1)_{\text{QG}}^{-1} \tilde{\omega} & \tilde{u} = \nabla^\perp \tilde{h} \end{cases}$$

[BC, SIMA 2013]

For **ill-prepared**  $U_0 \in H^m(\Omega)$  and  $t \lesssim 1$ ,

$$\|U^s - \tilde{U}\|_{H^{m-3}} \lesssim \varepsilon, \quad \left\| \int_0^t (U - \tilde{U}) \right\|_{H^{m-3}} \lesssim \varepsilon.$$



# Time-averaging and its connection to $\mathcal{P}$

- Time-averaging is used in observations and simulations.

(time-average  $\leftrightarrow$  time-integral),  $\partial_t U = \varepsilon^{-1} \mathcal{L}[U] + F \implies$   
$$\varepsilon(U(T) - U(0) - \int_0^T F) = \int_0^T \mathcal{L}[U] dt = \mathcal{L}[\int_0^T U^f].$$

## Bounded Inverse Theorem

Bounded, surjective, linear  $\mathcal{L} : X_1 \mapsto X_2$  ( $X_i$  being Banach), and  $\|(1 - \mathcal{P})U\|_{X_1} \approx \|U\|_{X_1/\ker\{\mathcal{L}\}}$  where  $\mathcal{P} = \text{proj onto } \ker\{\mathcal{L}\}$ . Then,

$$\left\| (1 - \mathcal{P})U \right\|_{X_1} \lesssim \left\| \mathcal{L}[U] \right\|_{X_2} \quad \text{i.e.} \quad \|U^f\|_{X_1} \lesssim \|\mathcal{L}[U^f]\|_{X_2}$$

Well-known examples:  $\mathcal{L} : \mathbf{R}^n \mapsto \mathbf{R}^m$  and  $\mathcal{L} = \text{div}$

**Time-averaging suppresses fast,** “  $\int_0^T \approx \mathcal{P}$  ”

$$\left\| \int_0^T U^f \right\|_{X_1} \lesssim \left\| \mathcal{L}[\int_0^T U^f] \right\|_{X_2} \lesssim \varepsilon.$$

# Time-averaging and transport equations

Idea useful in later slides, but also interesting on its own.

$u, \tilde{u}$ : velocities.       $s, \tilde{s}$ : (active) scalars so that  $u = \mathbf{q}[s]$ ,  $\tilde{u} = \mathbf{q}[\tilde{s}]$   
and  $\|u - \tilde{u}\|_{Y_1} \lesssim \|s - \tilde{s}\|_{Y_2}$ ,

$$\begin{aligned} \partial_t s + u \cdot \nabla s &= r, & s_0(\cdot) &= \tilde{s}_0(\cdot). \\ \partial_t \tilde{s} + \tilde{u} \cdot \nabla \tilde{s} &= \tilde{r}, \end{aligned}$$

Only assume,  $\int_0^t (r - \tilde{r}) \sim O(\varepsilon)$ ,

but allow,  $(r - \tilde{r}) \sim O(1)$  for  $t \in [0, O(1)]$ .

Write an eq for  $e := (s - \tilde{s}) + \int_0^t (r - \tilde{r}) d\tau$  and apply energy method

$$\|e(t)\| \lesssim \varepsilon, \quad \text{thus} \quad \|s(t) - \tilde{s}(t)\| \lesssim \varepsilon.$$

Results in relation to 2D RSW eq.

$$\begin{cases} \partial_t h + \nabla \cdot (hu) + \varepsilon^{-1} \nabla \cdot u = 0, \\ \partial_t u + u \cdot \nabla u + \varepsilon^{-1} u^\perp + \varepsilon^{-1} \nabla h = 0, \end{cases} \quad u \cdot \mathbf{n}|_{\partial\Omega} = 0.$$

**Well-posedness** of IBV problem of hyperbolic PDEs

- Linear.
  - $L^2$  theory: Friedrichs (58'), Lax & Phillips (60').
  - $H^m$  theory: Rauch & Massey (74', nonsingular boundary matrix), Rauch (85', singular boundary matrix, tangential regularity).
- Quasilinear.
  - Schochet (86'a, compressible Euler; 86'b, general case with **vorticity equation**) — full regularity.

**Singular limit** problem of multiscale Hyperbolic PDEs  
(Most results: compressible Euler  $\rightarrow$  incompressible Euler at zero-Mach-number limit)

- $\Omega = R^N$  or  $T^N$ 
  - well-prepared : Ebin, Kreiss, Beirao de Veiga, Klainerman-Majda, Tadmor (80s) ...
  - ill-prepared : Ukai, Isozaki(80s), Schochet, Masmoudi (90s) ...
  - geophysical models: Embid-Majda, Beale-Bourgeois, Babin-Mahalov-Nicolaenko (90s), Chemin-Desjardins-Gallagher-Grenier (book 06') ...
- Solid Wall on  $\partial\Omega$ 
  - Schochet (80s), Secchi, D. Jones (00s), ...
- Weak or strong convergence. No convergence rate.

- Link 1: extract slow dynamics.

- Apply  $\mathcal{P}$  on  $\partial_t U + B(U, U) = -\varepsilon^{-1} \mathcal{L}[U]$  and use  $\mathcal{P}\mathcal{L} \stackrel{?}{=} 0$

$$(\star) \quad \partial_t U^s = -\mathcal{P}B(U^s, U^s) - \mathcal{P}B(U^f, U^f) - \mathcal{P}B(U^s, U^f) - \mathcal{P}B(U^f, U^s)$$

- Slow-slow interaction, i.e.

$$\partial_t \tilde{U} = -\mathcal{P}B(\tilde{U}, \tilde{U}) \quad \text{with } \tilde{U}_0 \in \ker\{\mathcal{L}\}$$

$\stackrel{?}{\iff}$  The QG eq.

$$\partial_t \tilde{\omega} = -\nabla \cdot (\tilde{u} \tilde{\omega}), \quad \tilde{u} = \nabla^\perp (1 - \Delta)_{\text{QG}}^{-1} \tilde{\omega}$$

Note: need to show, with  $u \cdot \mathbf{n}|_{\partial\Omega} = 0$ , the ?'s are valid.

The nonlocal bound. cond. of  $(1 - \Delta)_{\text{QG}}^{-1}$  is crucial.

- Link 2: **time-averaging**.
  - For the Bounded Inverse Theorem to work, need to verify  $\mathcal{L}$  has closed graph (i.e. solvability). Or, directly use **elliptic estimates** from [S. Agmon, A. Douglis and L. Nirenberg, 1959 & 1964]

$$\text{With } u \cdot \mathbf{n} \Big|_{\partial\Omega} = 0, \quad \|U^f\|_{H^m(\Omega)} \lesssim \|\mathcal{L}[U^f]\|_{H^{m-1}(\Omega)}$$

$$\text{More or less, } \|u\|_{H^m} \lesssim (\|\nabla \times u\|_{H^{m-1}} + \|\nabla \cdot u\|_{H^{m-1}} + \|u\|_{L^2}).$$

- So,  $\int_0^T U^f \sim O(\varepsilon)$ . Also,  $(\star) \implies \partial_t U^s \sim O(1)$ .

$$\text{Integrating by parts } \implies \int_0^T \mathcal{P}B(U^f, U^s) \sim O(\varepsilon).$$

**Fast-fast** interaction  $\mathcal{P}B(U^f, U^f) \equiv 0$  by Calculus.

*Physical reason?*

- Link 3: Potential vorticity (PV)

$$\frac{\varepsilon^{-1} + \nabla \times u}{1 + \varepsilon h} = \varepsilon^{-1} + (\nabla \times u - h) + O(\varepsilon)$$

Define linear PV as  $\omega = \mathcal{K}[U] := (\nabla \times u - h)$  so that  $\mathcal{K}\mathcal{L} \equiv 0$  and

$$\partial_t \omega = -\nabla \cdot (u\omega) \quad \dots \quad (\text{eq of linear PV})$$

$\mathcal{L} = -\mathcal{L}^*$ ,  $\mathcal{P} = \mathcal{P}^* = \mathcal{K}^*(\mathcal{K}\mathcal{K}^*)_{\text{QG}}^{-1}\mathcal{K}$ , (least-square in frqncy space)

$\implies U^s = \mathcal{P}[U]$  and  $\omega = \mathcal{K}[U]$  are **1-to-1**.

$$\begin{aligned} \partial_t U^s &= -\mathcal{P}B(U^s, U^s) - \cancel{\mathcal{P}B(U^f, U^f)} \quad \longleftrightarrow \quad \partial_t \omega = -\nabla \cdot (u^s \omega) \\ &\quad -\mathcal{P}B(U^s, U^f) - \mathcal{P}B(U^f, U^s) \quad \quad \quad -\nabla \cdot (u^f \omega) \end{aligned}$$

- Final step: apply energy estimates to compare slow solution of

$$\partial_t \omega = -\nabla \cdot (u^s \omega) - \nabla \cdot (u^f \omega), \quad u^s = \nabla^\perp (1 - \Delta)_{\text{QG}}^{-1} \omega$$

and the QG solution of

$$\partial_t \tilde{\omega} = -\nabla \cdot (\tilde{u} \tilde{\omega}), \quad \tilde{u} = \nabla^\perp (1 - \Delta)_{\text{QG}}^{-1} \tilde{\omega}$$

knowing  $\int_0^T \nabla \cdot (u^f \omega) \sim O(\varepsilon)$

(Back to the **active scalar transport** eq!)



## Example 2: 2D fluid dynamics on a fast rotating sphere

### Incompressible Euler equations on $S^2$

$$\partial_t u + \nabla_u u + \nabla P = \frac{z}{\varepsilon} u^\perp, \quad \operatorname{div} u = 0$$

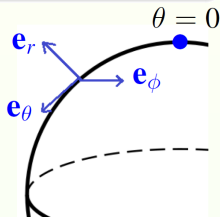
(admits Rossby-Haurwitz wave solutions)

$\theta$ : colatitude.       $\phi$ : longitude

$z = \cos \theta$ : **north-south variation** (a rugby will do)

$\ker\{\mathcal{L}\}$  are **zonal flows**

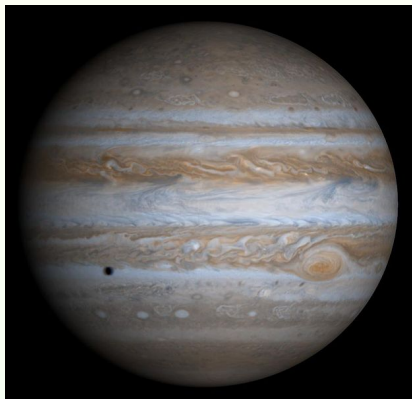
$\mathcal{P}u =$  zonal mean of  $u$



### Theorem [Cheng & Mahalov, E. J. Mech.-B/Fluids (2013)]

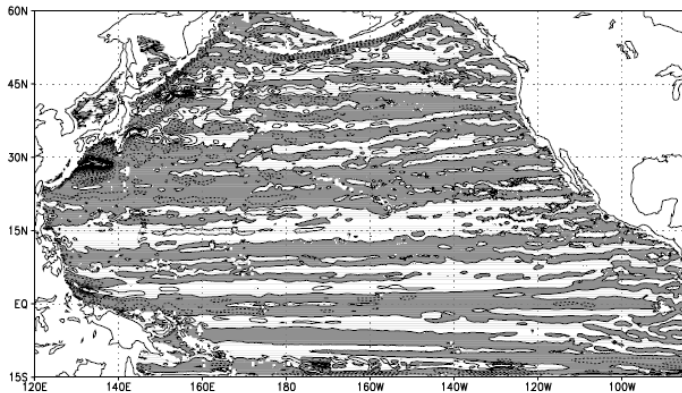
$$\left\| \int_0^T u(t, \cdot) dt - \int_0^T \mathcal{P}u(t, \cdot) dt \right\|_{H^{k-3}} \lesssim \varepsilon \quad \text{for } T \lesssim 1$$

- Proofs are coordinate free.
- Also, estimates/constants must be independent of  $\varepsilon$ .



**Figure :** True-color view of Jupiter, composed of 4 images taken by NASA's Cassini spacecraft on December 7, 2000. Credit: NASA/JPL/U. of Arizona.

## Zonal flows in simulations



Zonal jets at 1000 m depth in the North Pacific Ocean averaged over the last 5 years of a 58-year long computer simulation. Shaded and white areas are westward and eastward currents, respectively; the contour interval is  $2 \text{ cm s}^{-1}$ . [H. Nakano and H. Hasumi J. Phys. Oceanography (2005)]

### Viscous flows:

- [Lion-Temam-Wang 90s] primitive equations in shell domain.
- [Wirosoetisno 2013] Navier-Stokes eq. in rotating  $\mathbf{S}^2$ .  $\implies$  Global attractor is “almost” zonal, and is 0-dim for small  $\varepsilon$ .

### Inviscid flows:

- Incompressible Euler in rotating  $\mathbf{S}^2$ : governing **dynamics** of zonal flows  $\mathcal{P}u$ ? [Schochet 94' J. Diff. Eq.]: limiting dynamics *always*  $\exists$ .

PV eq. is 
$$\partial_t(\operatorname{curl} u - \frac{z}{\varepsilon}) + u \cdot \nabla(\operatorname{curl} u - \frac{z}{\varepsilon}) = 0$$

With **well**prepared initial data, limiting dynamics is **stationary**, zonal.

- Compressible shallow water equations on rotating  $\mathbf{S}^2$ . Even  $H^s$  estimates uniformly in  $Ro$ ,  $Fr$  are open problem.

$$\partial_t U = \frac{1}{\varepsilon} \mathcal{L}[U] + F$$

- Decompose  $U = U^s + U^f$  as well as the system.  $U^s \in \ker\{\mathcal{L}\}$ .
- Allow  $\partial_t U \sim O(\varepsilon^{-1})$ , namely, do not suppress fast  $U^f$ .
- Find vorticity operator  $\mathcal{K}$  s.t.  $\mathcal{K}\mathcal{L} = 0$  and establish correspondence between slow and **vortical dynamics**.
- Control of **fast-fast** interaction (?)
- **Verify**  $\|U^f\|_{X_1} \lesssim \|\mathcal{L}[U^f]\|_{X_2}$ , then,  $\int_0^T (1 - \mathcal{P})U dt \sim O(\varepsilon)$ .
- Time-averaging and flows.
- **Numerical analysis** – the importance to have schemes preserving vortical dynamics (i.e. slow dynamics)

Thank you!